GEOLOGIC SUMMARY SETTING

GEOLOGIC ATLAS OF THE MOON

RUMKER QUADRANGLE

I-805 (LAC 23)

The Rümker quadrangle, in the northwest quadrant of the Moon, is adjacent to the western rim of the multi-ring Imbrium basin and to Sinus Iridum, a large (220 km diameter) mare-filled crater. Both of these great depressions were probably formed by impact, as indicated here and elsewhere on the Moon by the characteristic form, distribution, and texture of surrounding materials and structures. The Imbrium basin and Iridum crater were filled by mare materials during the Imbrian and Eratosthenian Periods. In this quadrangle, the widespread ejecta blanket of the Imbrium basin, as well as the basin's concentric ridges and mountain rings, has largely been buried by terra materials of mixed origin and by ejecta from Iridum and numerous smaller impact craters. These materials, together with some terra units of probable volcanic origin, make up the highland terrain. Mare materials of Oceanus Procellarum cover part of this westwardsloping highland shelf and separate it from the Rümker Hills, an isolated plateau surrounded by the mare. No visible discontinuity distinguishes mare materials of Oceanus Procellarum from those in the Imbrium basin, both of which contain several units having similar albedo and color. Parts of the mare are very dark and smooth and appear relatively young, and parts of the terra contain many diverse landforms of Imbrian and younger age that appear to be volcanic. This region may therefore have been more active internally during the late stages of lunar history than many others on the near side.

GEOLOGIC UNITS

transection relations and, less definitively, by morphologic and other

criteria such as topographic sharpness, crater density, albedo, and

thermal response at eclipse. Where possible, all units are assigned to the

lunar stratigraphic column, as proposed by Shoemaker (1962), Shoe-

maker and Hackman (1962), and Wilhelms (1970). Materials are

grouped into five categories according to their physical characteristics.

probable mode of origin, or association with a major feature or event.

Imbrium basin materials

Imbrium basin materials comprise those units which by their age and

distribution appear to be more closely associated with the formation

of the Imbrium basin than with later volcanic, tectonic, or impact

events. Materials exposed in fault scarps bounding tilted blocks

marginal to the basin in the southeast highlands probably consist of

Imbrium ejecta and pre-basin bedrock. The units are indistinguishable

on the photographs and are mapped as one (IpIu). The Fra Mauro

Formation is questionably identified in the Rümker Hills, but else-

where the characteristic swirly texture of this Imbrium basin ejecta

blanket as well as its hilly counterpart, the Alpes Formation, is not

highland terrain around the Imbrium basin. These units resemble

subdued versions of both the Alpes and Fra Mauro Formations; thus

they are interpreted as mantling units of unknown composition, origin,

Iridum crater materials

the rim of Sinus Iridum just east of the quadrangle and provides a good

marker horizon for this region of the Moon. It covers Imbrium basin

materials and crater materials of early Imbrian age; superposition

relations of younger units indicate that it is of middle Imbrian age. As

with other large craters of impact origin, the rim materials of Iridum

are subdivided into members by their texture, depositional patterns,

and relief--characteristics which also reflect the size, mode of transport,

and velocity of the ejected or displaced materials. Within and beyond

the continuous ejecta blanket are numerous craters and clusters of

overlapping craters having elongate asymmetric outlines whose orien-

tation and disposition indicate secondary impacts generated by the

Iridum event. The largest individual members of the Iridum secondary

crater population are about 10 km in diameter. The ratio of maximum

secondary crater diameter to that of the primary crater is thus about

1:22, about the same as that around large craters like Copernicus or the

Mare and light plains materials

Two mare units of Eratosthenian and Imbrian age are distinguished

in Oceanus Procellarum by their albedo, color, and density of small

within the terra. Albedo and crater density apparently increase

progressively with increasing age: the youngest mare unit (Em) is

dark and smooth, whereas plains material (Ip) is light and highly pitted with small craters. Spectral reflectance (Whitaker, 1966) has also been useful for subdividing mare materials as color seems to be than younger ones, and the younger reflect more blue light. Although color may be related both to composition and to age, this concordance

between color and relative age may indicate a progressive change in

Imbrian craters more than 3 km across are much more numerous

on terrae than on maria, whereas Eratosthenian and Copernican

craters are about equally abundant on both types of terrain. The

Imbrian craters are embayed by both mare units but, with one

possible exception (Naumann G), Eratosthenian and younger craters

are superposed on the mare. These relations are in accord with a

latest Imbrian age for the oldest mare unit and a Copernican or late

Eratosthenian age for the mare embaying the Eratosthenian crater. The

paucity of partly buried craters or their outlines indicates that the

mare materials (probably basalts) are thick enough to have completely

covered most Imbrian craters and probably some pre-Imbrian craters

Crater materials

The age of the craters mapped here ranges from early Imbrian to

Copernican. A progressive decrease in the abundance of older to

younger craters on the terra probably reflects a declining flux of

impacting meteorites with time. Except for numerous secondary craters from Iridum, most craters have nearly circular outlines, rough

rims, and deep, flat to bowl-shaped floors characteristic of an impact

origin. Other craters have morphologies and material associations

Volcanic (?) landform materials Aside from the extensive basalt flows in Oceanus Procellarum, many other features in themselves or by association with others reflect the operation of volcanic processes in this region. The Gruithuisen domes (Ed) and some smaller similar domes mapped as the same unit resemble terrestrial cumulo-domes formed by extrusions of highly

viscous lava (Holmes, 1965, fig. 229). In the Rumker Hills, plug-like

forms which have arched and possibly penetrated to the surface are

bordered in places by lobate scarps resembling lava-flow fronts. Unlike

most upland areas, the Rümker Hills have a low albedo, comparable

with that of the adjacent mare, and, like the mare, they are probably

Within the mare in the southeast part of the quadrangle a series of

small craters (Ech) separated by mesa-like hills form a long, somewhat

sinuous chain resembling a string of beads. Most of these craters are

rimless, or nearly so (unlike impact craters) and resemble depressions

formed by caving along parts of a lava tube or by subsurface drainage

of magma into fissures and subsequent collapse. In this respect, the

crater chain probably represents an early and incomplete stage in the

process of rille formation; individual depressions did not coalesce and

develop into the winding, continuous, flat-floored valleys illustrated

Crater chains as well as several hills with summit craters or breached

ramparts (Ihc) in the western mare have morphologic analogs in

terrestrial volcanic fields (Scott and Trask, 1971). Craters with

(Trask and Titley, 1966). An outstanding terrestrial example of this

type of crater occurs in the Mount Lassen volcanic province (Chapman

and others, 1969, fig. 31-6). The floors of ringed craters as well as

some other craters in the Rümker quadrangle are filled with clusters of

small domes (Eldc). This association of clustered domes with craters resembling terrestrial volcanoes suggests that the domes may be tholoids consisting of viscous lava like those that form steep-sided domes in the vents of some strato-volcanoes on Earth. Other features, such as chain craters (ch), lobate material (I1), and a smooth unit (Es) on the rim of the crater Mairan, are also interpreted as volcanic

and suggest that magmatic differentiation progressed farther here than

STRUCTURE

The main rim of the Imbrium multi-ring basin is poorly defined

here. It originally formed high mountains like those of the Carpathian,

Apennine, and Caucasus around other parts of the basin. These have

subsided and been buried by mare basalts, Iridum crater ejecta, and

concentric with the Iridum crater than with the Imbrium basin. Some

ever, may be remnants of an Imbrium rim. Major mare ridges and rilles

(such as those of Rima Sharp) within Oceanus Procellarum strike

northwest to about latitude 40° 45° N., where their trends swing

rather abruptly to the northeast into Sinus Roris. This change in

direction roughly reflects the curvature of the Imbrium basin and

suggests that the ridges and rilles owe their existence (at least in

part) to faults and fractures concentric with the basin. The ridges

may represent upwelling and solidification of basaltic material along

fractures during late stages of mare extrusion, whereas the rilles, like some of the chain craters (Ech) previously discussed, are probably

Troughs and ridges radial to Sinus Iridum occur in the rim materials

of the Iridum crater and within terra units. Some of these lineaments

probably resulted from sculpturing of the terrain by missiles within the

ejecta cloud. However, many crater chains (ch) parallel lineament

the Iridum impact. It seems likely, therefore, that many lineaments, as

formed the Iridum crater eradicated much of the geologic record in the

Rümker quadrangle prior to middle Imbrian time. The Fra Mauro (?)

and Alpes Formations, which are characteristic of the earlier more

widespread ejecta blanket around the Imbrium basin, were mostly

buried by the Iridum crater materials as well as other younger units

Evidence for early Imbrian to Eratosthenian volcanism in the high-

lands is present (units Icfr, Ici, Elh, Eldc, Ed). (The more extensive

terra materials (Ith, Its) also may be partly of volcanic origin.) The

Rümker Hills were formed by faulting, uplift, and the intrusion of

penetrated it during the Imbrian and Eratosthenian Periods. During

the latter part of the Imbrian Period, lava flows filled crater floors and

depressions within the terra, forming the level plains (Ip) in this area.

These initial flows were succeeded by larger volumes of darker

materials, probably basalts, which filled great depressions and formed

the maria of Oceanus Procellarum, the Imbrium basin, and Iridum

crater. A change in composition of the basalt flows with time is

indicated by albedo and color variations. The later flows that probably

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large plug-like forms which may have locally arched the surface and

such as hilly and smooth terra and the mare materials of Oceanu

directions and, as they occur on plains deposits, must have formed after

GEOLOGIC HISTORY The extensive ejecta blanket produced by the large impact which

other materials. The large ridges of the Montes Jura are more nearly

of the high but subdued areas north of the Gruithuisen domes, how-

ringed floors (Icfr) are like those mapped elsewhere on the Moon

basaltic in composition or are mantled by basalt.

by Rima Sharp to the northwest.

elsewhere on the Moon.

collapsed lava drainage channels.

well as crater chains, have structural roots.

suggestive of an internal origin and are discussed below.

composition of the materials with time.

craters. Imbrian mare and plains materials fill small isolated depression

The ejecta blanket of the Iridum crater extends about 250 km from

and thickness.

Orientale basin.

recognized. Instead, two terra units (Ith, Its) make up most of the

The relative ages of units are determined by superposition and

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